Physical Insights from the Multidecadal Prediction of North Atlantic Sea Surface Temperature Variability Using Explainable Neural Networks

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Abstract

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Key Points:

\begin{itemize}
  \item Neural networks outperform persistence forecasts in predicting extreme states of North Atlantic sea surface temperature out to 25 years
  \item An explainable neural network technique reveals successful predictions rely consistently on the Transition Zone Region
  \item Predictions by neural networks trained on model output captures the phasing of multidecadal variability on an observation-based dataset
\end{itemize}

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North Atlantic sea surface temperatures, particularly in the subpolar region, are among the most predictable locations in the world’s oceans. However, it remains uncertain if processes in the atmosphere or ocean are more important for driving temperature fluctuations in this region occurring over multiple decades. We use a machine learning approach and train a neural network to predict the sea surface temperature state from climate model outputs, given snapshots of atmospheric or oceanic variables. Ocean variables lead to more accurate predictions relative to atmospheric variables and standard prediction baselines out to 25 years ahead if processes that drive the trends in climate, such as human-induced warming, are present in the data. These successful predictions arise consistently from the same region near the Gulf Stream-North Atlantic Current region. Despite being trained on climate models, the neural networks can predict the timing of observed positive and negative states of real-world sea surface temperatures, suggesting that there is potential for using model output to train neural networks at predicting the actual North Atlantic sea surface variability.

1 Introduction

Sea surface temperature (SST) anomalies averaged over the North Atlantic region exhibit alternating warm and cold periods on decadal timescales, known as the Atlantic Multidecadal Variability (AMV, or Atlantic Multidecadal Oscillation). The societal relevance of predicting AMV is underscored by linkages to multidecadal variations across multiple Earth system processes both within and beyond the North Atlantic (Zhang et al., 2019; Ruprich-Robert et al., 2021, and references therein). However, the dominant driver of AMV remains highly contested; leading contenders include ocean dynamics (Kim et al., 2018; Zhang et al., 2019; Arzel et al., 2022), atmospheric dynamics (Clement et al., 2015; Cane et al., 2017), and variations in external forcing (L. N. Murphy et al., 2021; Klavans et al., 2022). Each of these drivers imply different timescales of predictability, and the short observational record further complicates the disentanglement of their contributions.

Yet the subpolar North Atlantic (SPNA), the center of action for AMV, is considered among the most predictable locations for SST and ocean heat content across all ocean basins, with skill extending to decadal timescales (Buckley et al., 2019; Yeager, 2020). Mean wintertime mixed-layer depths reach over 1000 meters within the SPNA, resulting in large heat capacity that translates to long persistence and memory of SST anomalies (Deser et al., 2003; Holte et al., 2017). The SPNA encompasses key deep-water formation sites of the Atlantic Meridional Overturning Circulation (AMOC), and has been linked to multi-year to
multi-decadal predictability, both locally and in other regions such as the tropical Atlantic (Dunstone et al., 2011; Menary et al., 2015).

Current state-of-the-art approaches for decadal prediction of the climate system are often computationally intensive and highly sensitive to initial conditions, or constrained by assumptions of linearity in simplified models such as the Linear Inverse Model (Zanna, 2012; Huddart et al., 2017; Smith et al., 2019; Meehl et al., 2022). An alternative pathway emerges from neural networks (NN) and their ability to capture nonlinear processes and transformations (Hornik et al., 1989; Toms et al., 2020). NNs have successfully outperformed dynamical forecasts of El Niño-Southern Oscillation (ENSO) at interannual timescales (Ham et al., 2019) and detecting transitions between positive and negative states of the Pacific Decadal Oscillation (Gordon et al., 2021). Furthermore, recent developments of techniques such as Layer-wise Relevance Propagation (LRP) provide a way to peer into the “black box” of the NNs and identify the critical features for skillful predictions (Toms et al., 2020; Gordon et al., 2021; Wang et al., 2022). In this work, we investigate the potential of applying NNs to predicting NASST and use LRP to examine the relative importance of atmospheric and oceanic sources of predictability across multiple timescales.

2 Methods and Data

2.1 Datasets

We use the Community Earth System Model 1 (CESM1) Large Ensemble Simulations (LENS) based on a fully-coupled global climate model with nominal 1-degree resolution (Kay et al., 2015). We focus on a single model to investigate if NNs can learn the physics of NASST variability, without confounding factors and biases that arise from cross-model comparisons. CESM1 LENS features 42 members under the same external forcing but with slightly different atmospheric initial conditions, representing a comprehensive range of intrinsic climate variability. We use the historical period common across all ensemble members (1920 to 2005), totaling of 3,612 years of data for training, validation, and testing of the NNs.

To investigate if the predictability learned from CESM1 translates to a realistic dataset, we test the NNs on an observational dataset, the Hadley Center Sea Ice and Sea Surface Temperature (HadISST) that includes monthly data between 1870 and 2022 at 1-degree resolution (Rayner et al., 2003). Since the NNs require inputs of the same size, we re-grid HadISST to match the CESM1 resolution using bilinear interpolation.

2.2 Prediction Objective

The input features are 2-D annual mean snapshots of atmospheric and/or oceanic predictors (discussed in Section 2.3) over the North Atlantic (80 to 0°W, 0 to 65°N), and the output prediction is the state of NASST (either positive, negative, or neutral) a given number of years later (Fig. 1). The NASST index is the area-weighted, annual mean SST anomaly over the North Atlantic, essentially the unfiltered AMV Index (Ting et al., 2009). Considering recent work that suggests the importance of external forcing in driving AMV (L. N. Murphy et al., 2021; Klavans et al., 2022), we also examine differences in predictability of NASST with and without external forcings such as the anthropogenic warming trend, defined by the 42-member ensemble mean (referred to as forced and unforced, respectively).

We focus on predicting extreme NASST states due to its strong scientific and societal impacts. A 1-standard deviation (σ) threshold is used to separate the NASST into positive, negative, and neutral states (similar results are obtained using tercile thresholds). The threshold was selected to be high enough to distinguish extreme NASST anomalies, but low enough to permit sufficient samples for training. To prevent biases towards predicting a specific class simply due to its frequency of occurrence, following standard practice,
Figure 1. Schematic diagram of the NN prediction of NASST state using an example NASST-event in 1965 from ensemble member 37 of CESM1 LENS (Panel A). The snapshot of a selected predictor from 25 years prior (1940) is given to a FNN (Panel B), which outputs a prediction of the NASST state.

(Drummond et al., 2003; Buda et al., 2018; Gordon et al., 2021), we subsample the CESM1 output during training and validation so that there are equally 300 events per NASST state.

2.3 Atmospheric and Oceanic Predictors

To evaluate the importance of atmospheric versus oceanic drivers for NASST variability, we train networks to predict the NASST state given 2-D annual mean anomalies of the 4 following predictors:

1. **SST**, also used to calculate the NASST indices.
2. **Sea level pressure (SLP)**, an atmospheric predictor reflecting the state of the dominant atmospheric modes of variability in the region, e.g., the North Atlantic Oscillation (NAO)(Hurrell & Deser, 2010; Ruprich-Robert & Cassou, 2015).
3. **Sea surface salinity (SSS)**, an oceanic predictor that is not directly damped by heat fluxes to the atmosphere, allowing for the investigation of redistribution and damping by ocean circulation and its connections with NASST variability (Zhang, 2017).
4. **Sea surface height (SSH)**, an oceanic predictor used to infer geostrophic circulation with connections to variations in the strength of subpolar gyre (Koul et al., 2020). SSH is also related to subsurface ocean heat content with potential for long-term predictability (Buckley et al., 2019; Yeager, 2020).

These predictors are observable from the ocean surface, and are thus more likely to have longer records into the future with satellite observations, providing potential for application to operational predictions of climate. We tested additional predictors from CESM1, including net air-sea heat flux, barotropic streamfunction, mixed-layer depth, heat and salt...
content, and wind stress and its curl. None of these predictors yielded significantly better performance, so we focus on the above four variables.

Each predictor is cropped to the domain used to compute the NASST index. Ocean variables are re-gridded to match atmospheric grid using bilinear interpolation. We exclude regions over land and where the ice fraction exceeds 5% so that the NNs are given the same areas for each predictor. We normalize each predictor by dividing by 1σ across the time, space, and ensemble dimensions, ensuring comparable variability between predictors and equal numerical contribution during the training process (Singh & Singh, 2020). Multiple NNs are trained with each of the above mentioned predictors separately. NNs that include all predictors as input did not yield improved skill, but rather indicate equivalent accuracy to the best predictor at each leadtime (not shown).

2.4 Network Architecture and Training Procedure

To separately investigate the dependency in timescale and predictor, each NN is trained to predict the NASST state at a specific leadtime (t=0 to 25 years) given one predictor at a time. We withhold 10 members of CESM1 LENS for testing, and split the remaining 32 members into training (90%) and validation (10%) subsets. We initialize 100 different networks to account for randomness in the training process, totaling 10,400 networks (26 leadtimes × 4 predictors × 100 initialized networks). The training and validation sets are shuffled and resampled for each training iteration, ensuring that the results are not sensitive to a particular subset. Each network is trained for 50 epochs, but the training process is stopped if the validation loss increases for 5 consecutive epochs to prevent over-fitting. All discussed results are from the withheld testing set.

We explored combinations of architectures and hyperparameters for convolutional neural networks (CNNs) and fully-connected neural networks (FNNs). Both architectures yielded comparable performance (Fig. S1C). Our preliminary results with more complex networks did not produce significantly better results, but full exploration of other architectures is left for future work. Since our objective is not to tune network hyperparameters to maximize accuracy, but rather to gain physical insight on drivers of NASST variability by examining inter-predictor differences, we focus on the simpler FNN in this study containing 4 layers with 128 neurons each.

2.5 Prediction Baselines

We compare the accuracy of the trained NNs to two baselines. Since each class is evenly sampled during the training, there is a 33% chance that a given class will occur, which we set as the random chance baseline. We additionally examine the other extreme using the standard persistence baseline that assumes uninterrupted continuation of the current state (A. H. Murphy, 1992), which gives a stronger baseline than a damped persistence. For example, if the system is at NASST+ at the starting time (t=0 years), we assume it will also be NASST+ for the target leadtime.

3 Higher skill from oceanic predictors at multidecadal leadtimes in the presence of external forcing

We focus on the prediction skill for NASST+ and NASST- events (Fig. 2). For the predictions of Neutral events, the NNs had low accuracy equivalent to random chance. This is expected due to the challenge of predicting cases at the class boundaries or events with a weaker signal (Batista et al., 2004).

In the forced case (Fig. 2A-B), NNs outperform both persistence and random chance baselines regardless of the predictor. The atmospheric variable, SLP, has similar-to-worse accuracy at all leadtimes compared to SST. While this is unsurprising, considering the
short persistence timescales of the atmosphere in the extratropics, on the order of weeks (Frankignoul & Hasselmann, 1977), the NN still outperforms the persistence forecast and the random chance baseline for predicting NASST+ at all the leadtimes.

While SST appears to be a better predictor at earlier leadtimes, NNs trained by both oceanic predictors (SSS and SSH) achieve consistently higher accuracy than SST at decadal and longer leadtimes (Fig. 2A-B). Prolonged predictability from SSS could arise from absence of strong, direct damping by turbulent heat fluxes that exists in SSTs, allowing for more persistent SSS anomalies (Mignot & Frankignoul, 2003; Zhang, 2017). Similarly, subsurface heat content information present in SSH is shielded from damping by surface heat fluxes, leading to more persistence and potential predictability relative to SST (Deser et al., 2003; Buckley et al., 2019).

The increased predictability from oceanic variables is dependent upon the presence of external forcings. After removing the ensemble mean from the predictors and NASST index and repeating the training procedure, all NNs exhibit performance comparable to random chance after 5-10 years with minimal inter-predictor difference. This suggests both the importance of considering external forcing for climate prediction on multidecadal timescales and its differing impact on predictability derived from oceanic variables.

4 Consistent source of long-term predictability in the Transition Zone

We investigate the source of predictability for each predictor using LRP to examine the network’s decision-making process (Böhle et al., 2019). LRP back-propagates the “relevance” for given sample’s prediction from the final output node to the input layer of the NN. The total relevance is conserved during this process through a series of propagation rules, resulting in a “heatmap” of relevance indicating each pixel’s contribution to the network’s final decision (Montavon et al., 2019; Samek et al., 2021).
Figure 3. Composite relevance values (color) for "correct" NASST+ predictions of the top 50 performing networks for 0- to 25-year leadtimes, for the predictors from SST (A-E), SLP (F-J), SSS (K-O) and SSH (R-T), respectively. Relevance values are normalized for each composite. SSS relevance values were doubled to aid interpretability. Contours are the respective composites of standardized predictors for the given leadtime.
such relevance maps with known patterns of physical processes for predicting Pacific climate variability for possible correspondences (Toms et al., 2020; Gordon et al., 2021).

Since LRP produces the relevance map for a single sample, we examine the overall learned source of predictability by compositing relevances across correct predictions for the top 50 performing NNs of NASST+ and NASST-. The composites are normalized prior to visualization to have values between 0 and 1, though the raw output relevance is of order $10^{-4}$. We show relevance composites for key leadtimes between 0-25 years overlaid on composites of input predictors at corresponding leadtimes (Fig. 3) for the forced NASST+ cases. Results are broadly consistent in unforced and for NASST- cases (Fig. S2-S3).

For instantaneous predictions (leadtime 0), the relevance maps resemble known patterns associated with AMV and its drivers. For example the SST relevance map (Fig. 3E) captures the canonical horseshoe pattern of AMV (Zhang et al., 2019). Furthermore, the maximum relevance south of Newfoundland in SST, SSH, and SSS is collocated with the SPNA-Gulf Stream dipole associated with AMV-related SSTs and major ocean circulation features (Zhang, 2008; Nigam et al., 2018; Oelsmann et al., 2020; Gu & Gervais, 2022).

Interestingly, a second relevance maxima for SSS is present near the Amazon River outflow region, though further investigation is needed to determine if this is a model-dependent feature and its physical mechanisms. Overall, these aspects lend confidence that the NN has learned to rely upon regions that vary strongly with AMV and its associated ocean drivers.

Patterns associated with atmospheric drivers of NASST variability also emerge in relevance maps at leadtimes longer than 5 years (Fig. 3F-I). Successful predictions by SLP-trained NNs rely upon negative SLP anomalies near the Icelandic Low in the northeastern Atlantic, a center of action for NAO (Hurrell & Deser, 2010; Deser et al., 2010). This learned reliance on the NAO-NASST linkage without additional input is encouraging, suggesting that additional predictability beyond the persistence baseline achieved by SLP-trained NNs may arise from large-scale air-sea interaction in this region and resulting ocean circulation anomalies.

The Transition Zone between the subpolar and subtropical gyres emerges as a consistently important region for predicting NASST regardless of leadtime for oceanic predictors (Fig. 3K-T) (Buckley & Marshall, 2016). This region is influenced by AMOC and its associated fingerprint in surface and subsurface temperatures (Zhang, 2008). Relevance over this region remains high irrespective of the class (NASST+ or NASST-) or the presence of external forcing (Fig. S3). Since the NNs can derive multidecadal predictability of NASST by focusing on a region strongly influenced by AMOC, this result highlights the potential importance of ocean dynamics for determining the state of both forced and unforced NASST.

5 CESM1-trained neural networks predict the multidecadal oscillation of observed NASST states

Does the NNs' skill for NASST prediction apply beyond the CESM1 model world? Because of the limited observational record of SSH, SSS, and SLP, we test if NNs trained on CESM1 SSTs can successfully predict the NASST state in HadISST. Accounting for reductions due to the 25 year leadtime, there remains 128 years of data between 1895 to 2022. The $1\sigma$ threshold ($0.55^\circ C$) yielded 29 (17) NASST+ (NASST-) events. The distribution is skewed due to the warming trend. Due to the limited samples, we do not perform transfer learning for the HadISST dataset and the accuracy values were noisy, particularly at long leadtimes. Therefore, we focus broadly on the frequency of predictions by class (Fig. 4).

The frequency of predictions by class across all NNs aligns with the multidecadal oscillation of the NASST in HadISST, including larger frequency of NASST- pre-1925, between 1960-1990, and the intervening warm periods. This is true particularly for interannual and
Figure 4. Frequency of predicted class of each target year aggregated for interannual (1-9 years) (A), decadal (10-19 years) (B), and multidecadal (20-25 years) (C) lead times for the HadISST (in colored bars). Blue/red/gray bars are the frequency of the negative/positive/neutral NASST predictions. The NASST Index from HadISST (solid-black line) and 1σ thresholds (dashed-black lines) are shown for reference.
multidecadal leadtimes (Fig. 4A,C), with shifted phasing at decadal leadtimes (Fig. 4B).
The same results are recovered for the unforced case, though the multidecadal phasing of
predictions is nearly absent for the decadal leadtimes (Fig. S4). These are surprising results
for two main reasons: The first is that the NN is not simply predicting the anthropogenic
warming trend (e.g. monotonically increasing NASST+ predictions in time), but instead
has successfully learned the non-linear, oscillatory behavior of the observed NASST index.
The second is that the weights not have been re-adjusted to HadISST, revealing that NNs
trained on potentially biased CESM1 output maintain their ability to predict the phasing
of observed multidecadal climate variability. Overall, this suggests promise for applying
NNs trained on model output to predicting the general details and trajectory of non-linear
multidecadal climate variability in corresponding observational datasets such as HadISST.

6 Discussion and Summary

We investigated the potential of applying NNs to multidecadal prediction of NASST
variability and using LRP to understand the relative contributions of oceanic and atmos-
pheric drivers. Three main conclusions of this work are:

1. NNs trained with oceanic variables can predict NASST+ and NASST- states on
   multidecadal timescales, outperforming persistence and random chance baselines in
   the presence of external forcing.
2. The Transition Zone emerges as consistent region from where NNs derive predictive
   skill, regardless of prediction leadtime, NASST state, and the presence of the external
   forcing, suggesting a connection to ocean dynamics such as AMOC.
3. NNs trained on CESM1 were able to predict the multidecadal phasing of observed
   NASST states without weight readjustment, suggesting promise for training NNs
   using model output for multidecadal prediction of observed climate.

While increased predictive skill from oceanic variables highlights the importance of
ocean dynamics for multidecadal NASST variability, we find that this depends upon the
presence of external forcing. There is little difference in skill between the predictors in
unforced case, suggesting that external forcing differently impacts predictability derived
from oceanic and atmospheric variables. A possible explanation is the larger heat capacity
of the ocean allows for the integration of the externally forced signal, leading to increased
predictability on multidecadal and longer timescales (Frankignoul & Hasselmann, 1977).

The high-relevance over the Transition Zone region is remarkably consistent across
timescales in both unforced and forced cases. This region corresponds to the maximum
loading in the AMOC fingerprint, suggesting that the dynamics driving both internal and
external NASST variability are collocated and linked to ocean dynamics (Zhang, 2008).
Predictability arising from a stationary feature in a single region, rather than smaller-
scale features that propagate across the domain, might also explain why the simpler FNN
performed comparably to CNNs; For predicting NASST, the absolute position of the feature
is more important than its translation invariance, erasing the advantage conferred by the
CNN architecture (Barnes et al., 2022).

A cautionary note is that higher accuracy from networks trained with oceanic predictors
could be a model dependent feature. Our results are focused on NNs trained with CESM1,
a coarse-resolution model with biases in the separation of the Gulf Stream and position of
the North Atlantic Current (Kirtman et al., 2012). Since our relevance maps reveal that
NNs depend upon this region for skillful predictions of NASST state, verifying the model
dependence of this aspect by training NNs with other model large ensembles, reanalyses, or
observational datasets is an important future endeavor. Considering connections between
biases in mean state and decadal variability over the SPNA, exploring correspondences
between the resultant relevance maps and biases in ocean circulation may unveil further hints on the importance of ocean dynamics for NASST predictability (Menary et al., 2015).

**Open Research Section**

Datasets for this research are available in these in-text data citation references: (Kay et al., 2015), (Rayner et al., 2003). The monthly output from the CESM1 Large Ensemble is publicly available from the National Center for Atmospheric Research’s Climate Data Gateway on the Earth System Grid (https://www.cesm.ucar.edu/community-projects/lens/data-sets/). Monthly variables TS, LANDFRAC, ICEFRAC, SSS, PSL, and SSH were used for this study, and further specific instructions on accessing the output for CESM1 is detailed at this link: (https://www.cesm.ucar.edu/community-projects/lens/data-sets). The HadISST dataset can be downloaded directly from their website (https://www.metoffice.gov.uk/hadobs/hadisst/).

Software for this work is available on Zenodo (DOI: https://doi.org/10.5281/zenodo.8342739), and the corresponding linked github repository (https://github.com/glennliu265/predict_nasst). The data will be The Pytorch-LRP Software is available from the in-text data citation reference: (Böhle et al., 2019) and can be found in the following repository (https://github.com/moboehle/Pytorch-LRP).

**Acknowledgments**

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Figure 2.
Figure 4.
Physical Insights from the Multidecadal Prediction of
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To investigate if the predictability learned from CESM1 translates to a realistic dataset, we test the NNs on an observational dataset, the Hadley Center Sea Ice and Sea Surface Temperature (HadISST) that includes monthly data between 1870 and 2022 at 1-degree resolution (Rayner et al., 2003). Since the NNs require inputs of the same size, we re-grid HadISST to match the CESM1 resolution using bilinear interpolation.

2.2 Prediction Objective

The input features are 2-D annual mean snapshots of atmospheric and/or oceanic predictors (discussed in Section 2.3) over the North Atlantic (80 to 0°W, 0 to 65°N), and the output prediction is the state of NASST (either positive, negative, or neutral) a given number of years later (Fig. 1). The NASST index is the area-weighted, annual mean SST anomaly over the North Atlantic, essentially the unfiltered AMV Index (Ting et al., 2009). Considering recent work that suggests the importance of external forcing in driving AMV (L. N. Murphy et al., 2021; Klavans et al., 2022), we also examine differences in predictability of NASST with and without external forcings such as the anthropogenic warming trend, defined by the 42-member ensemble mean (referred to as forced and unforced, respectively).

We focus on predicting extreme NASST states due to its strong scientific and societal impacts. A 1-standard deviation (σ) threshold is used to separate the NASST into positive, negative, and neutral states (similar results are obtained using tercile thresholds). The threshold was selected to be high enough to distinguish extreme NASST anomalies, but low enough to permit sufficient samples for training. To prevent biases towards predicting a specific class simply due to its frequency of occurrence, following standard practice
Figure 1. Schematic diagram of the NN prediction of NASST state using an example NASST-event in 1965 from ensemble member 37 of CESM1 LENS (Panel A). The snapshot of a selected predictor from 25 years prior (1940) is given to a FNN (Panel B), which outputs a prediction of the NASST state.

(Drummond et al., 2003; Buda et al., 2018; Gordon et al., 2021), we subsample the CESM1 output during training and validation so that there are equally 300 events per NASST state.

2.3 Atmospheric and Oceanic Predictors

To evaluate the importance of atmospheric versus oceanic drivers for NASST variability, we train networks to predict the NASST state given 2-D annual mean anomalies of the 4 following predictors:

1. **SST**, also used to calculate the NASST indices.
2. **Sea level pressure (SLP)**, an atmospheric predictor reflecting the state of the dominant atmospheric modes of variability in the region, e.g., the North Atlantic Oscillation (NAO) (Hurrell & Deser, 2010; Ruprich-Robert & Cassou, 2015).
3. **Sea surface salinity (SSS)**, an oceanic predictor that is not directly damped by heat fluxes to the atmosphere, allowing for the investigation of redistribution and damping by ocean circulation and its connections with NASST variability (Zhang, 2017).
4. **Sea surface height (SSH)**, an oceanic predictor used to infer geostrophic circulation with connections to variations in the strength of subpolar gyre (Koul et al., 2020). SSH is also related to subsurface ocean heat content with potential for long-term predictability (Buckley et al., 2019; Yeager, 2020).

These predictors are observable from the ocean surface, and are thus more likely to have longer records into the future with satellite observations, providing potential for application to operational predictions of climate. We tested additional predictors from CESM1, including net air-sea heat flux, barotropic streamfunction, mixed-layer depth, heat and salt.
content, and wind stress and its curl. None of these predictors yielded significantly better
performance, so we focus on the above four variables.

Each predictor is cropped to the domain used to compute the NASST index. Ocean
variables are re-gridded to match atmospheric grid using bilinear interpolation. We exclude
regions over land and where the ice fraction exceeds 5% so that the NNs are given the same
areas for each predictor. We normalize each predictor by dividing by $1\sigma$ across the time,
space, and ensemble dimensions, ensuring comparable variability between predictors and
equal numerical contribution during the training process (Singh & Singh, 2020). Multiple
NNs are trained with each of the above mentioned predictors separately. NNs that include
all predictors as input did not yield improved skill, but rather indicate equivalent accuracy
to the best predictor at each leadtime (not shown).

2.4 Network Architecture and Training Procedure

To separately investigate the dependency in timescale and predictor, each NN is trained
to predict the NASST state at a specific leadtime ($t=0$ to 25 years) given one predictor at
a time. We withhold 10 members of CESM1 LENS for testing, and split the remaining
32 members into training (90%) and validation (10%) subsets. We initialize 100 different
networks to account for randomness in the training process, totaling 10,400 networks (26
leadtimes $\times$ 4 predictors $\times$ 100 initialized networks). The training and validation sets are
shuffled and resampled for each training iteration, ensuring that the results are not sensitive
to a particular subset. Each network is trained for 50 epochs, but the training process is
stopped if the validation loss increases for 5 consecutive epochs to prevent over-fitting. All
discussed results are from the withheld testing set.

We explored combinations of architectures and hyperparameters for convolutional neu-
networks (CNNs) and fully-connected neural networks (FNNs). Both architectures
yielded comparable performance (Fig. S1C). Our preliminary results with more complex
networks did not produce significantly better results, but full exploration of other architec-
tures is left for future work. Since our objective is not to tune network hyperparameters to
maximize accuracy, but rather to gain physical insight on drivers of NASST variability by
examining inter-predictor differences, we focus on the simpler FNN in this study containing
4 layers with 128 neurons each.

2.5 Prediction Baselines

We compare the accuracy of the trained NNs to two baselines. Since each class is evenly
sampled during the training, there is a 33% chance that a given class will occur, which we
set as the random chance baseline. We additionally examine the other extreme using the
standard persistence baseline that assumes uninterrupted continuation of the current state
(A. H. Murphy, 1992), which gives a stronger baseline than a damped persistence. For
example, if the system is at NASST+ at the starting time ($t=0$ years), we assume it will
also be NASST+ for the target leadtime.

3 Higher skill from oceanic predictors at multidecadal leadtimes in the
presence of external forcing

We focus on the prediction skill for NASST+ and NASST- events (Fig. 2). For the
predictions of Neutral events, the NNs had low accuracy equivalent to random chance. This
is expected due to the challenge of predicting cases at the class boundaries or events with a
weaker signal (Batista et al., 2004).

In the forced case (Fig. 2A-B), NNs outperform both persistence and random chance
baselines regardless of the predictor. The atmospheric variable, SLP, has similar-to-worse
accuracy at all leadtimes compared to SST. While this is unsurprising, considering the
Figure 2. The mean accuracy by leadtime for predicting NASST+ and NASST- states for NNs trained with each predictor. X-axis is the prediction leadtime from 0 to 25 years. Shading indicates the 95% standard error of 100 NNs for each predictor. NNs trained with oceanic predictors SSH (blue) and SSS (pink) outperform those trained with SST (red) and SLP (yellow) at long leadtimes in the forced case (A-B). For the unforced case (C-D), performance is similar to the random chance baselines after 5-10 years (C-D).

short persistence timescales of the atmosphere in the extratropics, on the order of weeks (Frankignoul & Hasselmann, 1977), the NN still outperforms the persistence forecast and the random chance baseline for predicting NASST+ at all the leadtimes.

While SST appears to be a better predictor at earlier leadtimes, NNs trained by both oceanic predictors (SSS and SSH) achieve consistently higher accuracy than SST at decadal and longer leadtimes (Fig. 2A-B). Prolonged predictability from SSS could arise from absence of strong, direct damping by turbulent heat fluxes that exists in SSTs, allowing for more persistent SSS anomalies (Mignot & Frankignoul, 2003; Zhang, 2017). Similarly, subsurface heat content information present in SSH is shielded from damping by surface heat fluxes, leading to more persistence and potential predictability relative to SST (Deser et al., 2003; Buckley et al., 2019).

The increased predictability from oceanic variables is dependent upon the presence of external forcings. After removing the ensemble mean from the predictors and NASST index and repeating the training procedure, all NNs exhibit performance comparable to random chance after 5-10 years with minimal inter-predictor difference. This suggests both the importance of considering external forcing for climate prediction on multidecadal timescales and its differing impact on predictability derived from oceanic variables.

4 Consistent source of long-term predictability in the Transition Zone

We investigate the source of predictability for each predictor using LRP to examine the network’s decision-making process (Böhle et al., 2019). LRP back-propagates the “relevance” for given sample’s prediction from the final output node to the input layer of the NN. The total relevance is conserved during this process through a series of propagation rules, resulting in a “heatmap” of relevance indicating each pixel’s contribution to the network’s final decision (Montavon et al., 2019; Samek et al., 2021). Previous works compared
Figure 3. Composite relevance values (color) for "correct" NASST+ predictions of the top 50 performing networks for 0- to 25-year leadtimes, for the predictors from SST (A-E), SLP (F-J), SSS (K-O) and SSH (R-T), respectively. Relevance values are normalized for each composite. SSS relevance values were doubled to aid interpretability. Contours are the respective composites of standardized predictors for the given leadtime.
such relevance maps with known patterns of physical processes for predicting Pacific climate variability for possible correspondences (Toms et al., 2020; Gordon et al., 2021).

Since LRP produces the relevance map for a single sample, we examine the overall learned source of predictability by compositing relevances across correct predictions for the top 50 performing NNs of NASST+ and NASST-. The composites are normalized prior to visualization to have values between 0 and 1, though the raw output relevance is of order $10^{-4}$. We show relevance composites for key leadtimes between 0-25 years overlaid on composites of input predictors at corresponding leadtimes (Fig. 3) for the forced NASST+ cases. Results are broadly consistent in unforced and for NASST- cases (Fig. S2-S3).

For instantaneous predictions (leadtime 0), the relevance maps resemble known patterns associated with AMV and its drivers. For example the SST relevance map (Fig. 3E) captures the canonical horseshoe pattern of AMV (Zhang et al., 2019). Furthermore, the maximum relevance south of Newfoundland in SST, SSH, and SSS is collocated with the SPNA-Gulf Stream dipole associated with AMV-related SSTs and major ocean circulation features (Zhang, 2008; Nigam et al., 2018; Oelsmann et al., 2020; Gu & Gervais, 2022). Interestingly, a second relevance maxima for SSS is present near the Amazon River outflow region, though further investigation is needed to determine if this is a model-dependent feature and its physical mechanisms. Overall, these aspects lend confidence that the NN has learned to rely upon regions that vary strongly with AMV and its associated ocean drivers.

Patterns associated with atmospheric drivers of NASST variability also emerge in relevance maps at leadtimes longer than 5 years (Fig. 3F-I). Successful predictions by SLP-trained NNs rely upon negative SLP anomalies near the Icelandic Low in the northeastern Atlantic, a center of action for NAO (Hurrell & Deser, 2010; Deser et al., 2010). This learned reliance on the NAO-NASST linkage without additional input is encouraging, suggesting that additional predictability beyond the persistence baseline achieved by SLP-trained NNs may arise from large-scale air-sea interaction in this region and resulting ocean circulation anomalies.

The Transition Zone between the subpolar and subtropical gyres emerges as a consistently important region for predicting NASST regardless of leadtime for oceanic predictors (Fig. 3K-T) (Buckley & Marshall, 2016). This region is influenced by AMOC and its associated fingerprint in surface and subsurface temperatures (Zhang, 2008). Relevance over this region remains high irrespective of the class (NASST+ or NASST-) or the presence of external forcing (Fig. S3). Since the NNs can derive multidecadal predictability of NASST by focusing on a region strongly influenced by AMOC, this result highlights the potential importance of ocean dynamics for determining the state of both forced and unforced NASST.

5 CESM1-trained neural networks predict the multidecadal oscillation of observed NASST states

Does the NNs’ skill for NASST prediction apply beyond the CESM1 model world? Because of the limited observational record of SSH, SSS, and SLP, we test if NNs trained on CESM1 SSTs can successfully predict the NASST state in HadISST. Accounting for reductions due to the 25 year leadtime, there remains 128 years of data between 1895 to 2022. The 1σ threshold (0.55°C) yielded 29 (17) NASST+ (NASST-) events. The distribution is skewed due to the warming trend. Due to the limited samples, we do not perform transfer learning for the HadISST dataset and the accuracy values were noisy, particularly at long leadtimes. Therefore, we focus broadly on the frequency of predictions by class (Fig. 4).

The frequency of predictions by class across all NNs aligns with the multidecadal oscillation of the NASST in HadISST, including larger frequency of NASST- pre-1925, between 1960-1990, and the intervening warm periods. This is true particularly for interannual and
Figure 4. Frequency of predicted class of each target year aggregated for interannual (1-9 years) (A), decadal (10-19 years) (B), and multidecadal (20-25 years) (C) lead times for the HadISST (in colored bars). Blue/red/gray bars are the frequency of the negative/positive/neutral NASST predictions. The NASST Index from HadISST (solid-black line) and 1σ thresholds (dashed-black lines) are shown for reference.
multidecadal leadtimes (Fig. 4A,C), with shifted phasing at decadal leadtimes (Fig. 4B).

The same results are recovered for the unforced case, though the multidecadal phasing of predictions is nearly absent for the decadal leadtimes (Fig. S4). These are surprising results for two main reasons: The first is that the NN is not simply predicting the anthropogenic warming trend (e.g., monotonically increasing NASST+ predictions in time), but instead has successfully learned the non-linear, oscillatory behavior of the observed NASST index. The second is that the weights not have been re-adjusted to HadISST, revealing that NNs trained on potentially biased CESM1 output maintain their ability to predict the phasing of observed multidecadal climate variability. Overall, this suggests promise for applying NNs trained on model output to predicting the general details and trajectory of non-linear multidecadal climate variability in corresponding observational datasets such as HadISST.

6 Discussion and Summary

We investigated the potential of applying NNs to multidecadal prediction of NASST variability and using LRP to understand the relative contributions of oceanic and atmospheric drivers. Three main conclusions of this work are:

1. NNs trained with oceanic variables can predict NASST+ and NASST- states on multidecadal timescales, outperforming persistence and random chance baselines in the presence of external forcing.

2. The Transition Zone emerges as consistent region from where NNs derive predictive skill, regardless of prediction leadtime, NASST state, and the presence of the external forcing, suggesting a connection to ocean dynamics such as AMOC.

3. NNs trained on CESM1 were able to predict the multidecadal phasing of observed NASST states without weight readjustment, suggesting promise for training NNs using model output for multidecadal prediction of observed climate.

While increased predictive skill from oceanic variables highlights the importance of ocean dynamics for multidecadal NASST variability, we find that this depends upon the presence of external forcing. There is little difference in skill between the predictors in unforced case, suggesting that external forcing differently impacts predictability derived from oceanic and atmospheric variables. A possible explanation is the larger heat capacity of the ocean allows for the integration of the externally forced signal, leading to increased predictability on multidecadal and longer timescales (Frankignoul & Hasselmann, 1977).

The high-relevance over the Transition Zone region is remarkably consistent across timescales in both unforced and forced cases. This region corresponds to the maximum loading in the AMOC fingerprint, suggesting that the dynamics driving both internal and external NASST variability are colocated and linked to ocean dynamics (Zhang, 2008). Predictability arising from a stationary feature in a single region, rather than smaller-scale features that propagate across the domain, might also explain why the simpler FNN performed comparably to CNNs; For predicting NASST, the absolute position of the feature is more important than its translation invariance, erasing the advantage conferred by the CNN architecture (Barnes et al., 2022).

A cautionary note is that higher accuracy from networks trained with oceanic predictors could be a model dependent feature. Our results are focused on NNs trained with CESM1, a coarse-resolution model with biases in the separation of the Gulf Stream and position of the North Atlantic Current (Kirtman et al., 2012). Since our relevance maps reveal that NNs depend upon this region for skillful predictions of NASST state, verifying the model dependence of this aspect by training NNs with other model large ensembles, reanalyses, or observational datasets is an important future endeavor. Considering connections between biases in mean state and decadal variability over the SPNA, exploring correspondences
between the resultant relevance maps and biases in ocean circulation may unveil further hints on the importance of ocean dynamics for NASST predictability (Menary et al., 2015).

Open Research Section

Datasets for this research are available in these in-text data citation references: (Kay et al., 2015), (Rayner et al., 2003). The monthly output from the CESM1 Large Ensemble is publicly available from the National Center for Atmospheric Research’s Climate Data Gateway on the Earth System Grid (https://www.cesm.ucar.edu/community-projects/lens/data-sets/). Monthly variables TS, LANDFRAC, ICEFRAC, SSS, PSL, and SSH were used for this study, and further specific instructions on accessing the output for CESM1 is detailed at this link: (https://www.cesm.ucar.edu/community-projects/lens/data-sets). The HadISST dataset can be downloaded directly from their website (https://www.metoffice.gov.uk/hadobs/hadisst/).

Software for this work is available on Zenodo (DOI: https://doi.org/10.5281/zenodo.8342739), and the corresponding linked github repository (https://github.com/glennliu265/predict_nasst). The data will be The Pytorch-LRP Software is available in from the in-text data citation reference: (Böhle et al., 2019) and can be found in the following repository (https://github.com/moboehle/Pytorch-LRP).

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Supporting Information for "Physical Insights from the Multidecadal Prediction of North Atlantic Sea Surface Temperature Variability Using Explainable Neural Networks"

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Introduction

For the supporting information, we provide details on the hyperparameters of the fully-connected neural network (FNN) used in this project (Table S1). We compare the performance between FNNs and convolutional neural networks (CNNs) (Fig. S1). Additional figures are also provided for different cases discussed in the main text. They demonstrate that the main conclusions are not sensitive to these different cases.
Table S1. Neural Network Architecture and Training Hyperparameters used in the Fully-Connected Neural Network (FNN)

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<th>Value</th>
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<td>Number of layers</td>
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<td>Neurons per layer</td>
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<td>Activation Function</td>
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<td>Dropout Percentage*</td>
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<td>Learning Rate</td>
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*Dropout layer included prior the the last layer.
Figure S1. Schematic diagram of an example AMV Prediction problem (A) for the 2-layer CNN (B). The comparison in positive and negative North Atlantic Sea Surface Temperature (NASST+, NASST-) test accuracy between the FNN (yellow) and CNN (blue) for an SST predictor (C), with the random chance (dotted) and persistence baselines (black). Both networks perform similarly regardless of predictor, and their means are largely within the 95% standard error across initialized networks (shading).
Figure S2. Same as Figure 3, but for “correct” NASST- predictions for the top 50 performing networks. The regions of high relevance, i.e., sources of predictability, resemble that of NASST+, though there are small differences. The AMV maximum in the central subpolar gyre is more distinctly outline for SST at leadtime 0 (Panel E). Additionally, the NN focuses on anomalies closer to the Azores High at 5-year leadtimes, rather than directly to the Iceland low as in the NASST+ case (Panel I).
Figure S3. Same as Fig. 3, but for the unforced case where the external forcing was removed. The regions of maximum relevance resemble that of the forced NASST+ predictions. This similarity between both forced and unforced cases suggests that the NASST predictability is sourced from similar regions and dynamics, though further work is needed to explicitly investigate the responsible processes.
Figure S4. Same as Figure 4, but for NNs trained with unforced CESM1 data predicting the NASST Index from HadISST, detrended by removing a cubic fit. The result is not sensitive to the detrending method.