The 2020 COVID-19 pandemic and atmospheric composition: back to the future

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Abstract

The COVID-19 global pandemic and associated government lockdowns dramatically altered human activity, providing a window into how changes in individual behavior, enacted en masse, impact atmospheric composition. The resulting reductions in anthropogenic activity represent an unprecedented event that yields a glimpse into both the past and a future where emissions to the atmosphere are reduced. While air pollutants and greenhouse gases share many common anthropogenic sources, there is a sharp difference in the response of their atmospheric concentrations to COVID-19 emissions changes due in large part to their different lifetimes. Here, we discuss the lessons learned from the COVID-19 disruptions for future mitigation strategies and our current and future Earth observing system.
Societal shifts due to COVID-19 reveal large-scale complexities and feedbacks between atmospheric chemistry and climate change


The COVID-19 global pandemic and associated lockdowns dramatically altered human activity, providing a window into how changes in individual behavior, enacted en masse, impact atmospheric composition. The resulting reductions in anthropogenic activity represent an unprecedented event that yields a glimpse into a future where emissions to the atmosphere are reduced. While air pollutants and greenhouse gases share many common anthropogenic sources, there is a sharp difference in the response of their atmospheric concentrations to COVID-19 emissions changes due in large part to their different lifetimes. Here, we discuss two key takeaways from modeling and observational studies. First, despite dramatic declines in mobility and associated vehicular emissions, the atmospheric growth rates of greenhouse gases were not slowed. Second, it demonstrated empirically that the response of atmospheric composition to emissions changes is heavily modulated by factors including carbon cycle feedbacks to CH₄ and CO₂, background pollutant levels, the timing and location of emissions changes, and climate feedbacks on air quality.

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The authors declare no competing interests.

JLL led the manuscript and the human activity analysis. JN, DS, and POW led the study team. K. Barsanti, K. Bowman, DS, AF, and EK led study subgroups. Remaining authors contributed data analysis. All authors helped revise the manuscript.

Significance Statement

The COVID-19 pandemic and associated lockdowns caused significant changes to human activity that temporarily altered our imprint on the atmosphere, providing a brief glimpse of potential future changes in atmospheric composition. This event showed key differences in how air quality and atmospheric greenhouse gas concentrations respond to changes in anthropogenic emissions, with implications for future mitigation strategies.
in anthropogenic emissions during 2020. Second, we examine how the reduction in CO$_2$ emissions impacted the atmospheric CO$_2$ growth rate. Third, we show that the response of AQ to emissions reductions is very spatially heterogeneous, and summarize the causes of that heterogeneity. Fourth, we discuss the implications of these results for future AQ improvement strategies, our understanding of processes controlling GHG concentrations in the atmosphere, feedbacks between AQ, GHGs, and climate, and finally close by identifying strengths and gaps in our current observing networks. We draw three primary conclusions from this synthesis:

1. Despite drastic reductions in mobility and resulting vehicular emissions during 2020, the growth rates of GHGs in the atmosphere were not slowed.

2. The lack of clear declines in the atmospheric growth rates of CO$_2$ and CH$_4$, despite large reductions in human activity, reflect carbon cycle feedbacks in air-sea carbon exchange, large interannual variability in the land carbon sink, and the chemical lifetime of CH$_4$. These feedbacks foreshadow similar challenges to intentional mitigation.

3. The response of AQ to emissions changes is heavily modulated by factors including background pollutant levels, the timing and location of emissions changes, and climate-related factors like heat waves and wildfires. Achieving robust improvements to AQ thus require sustained reductions of both AQ and GHG emissions.

**Summary of emissions in 2020**

As AQ-relevant gases and CO$_2$ are co-emitted by combustion processes, decreases in human activity are expected to drive decreases in both of these species. Figure 2 summarizes changes to key sectors of human activity during the COVID-19 pandemic. Figure 2a shows the Oxford Stringency Index (1), which quantifies the severity of government-imposed restrictions on travel, businesses, schools, and other aspects of society. Panels b, c, and d show changes in air travel & maritime shipping, traffic, and United States (US) electricity use, respectively. There is a clear decrease in air travel and traffic for most of the world in March 2020, when the first major wave of COVID-19 led governments to institute quarantine measures (see also high values of the Stringency Index). Maritime shipping (to west coast US ports) and power generation (in the US) were less affected. Power generation in particular remained within approximately 5% of 2019 levels.

Reductions in NO$_x$ emissions were apparent in both in situ (5) and satellite (6) observations of NO$_2$ concentrations due to the short atmospheric lifetime of NO$_x$ (\(< 1\) day). Estimates of NO$_x$ emissions reductions from assimilating satellite data in global models (7), combining global chemical models with machine learning trained on surface measurements (8), or activity data (including electricity use, traffic/mobility data, flight data, etc.) (9–11) find regional reductions of 10% to 40% during the strictest lockdown periods. Generally, methods assimilating satellite data report smaller reductions (10% to 20%) than studies based on activity data (25% to 40%).
Fig. 2. Metrics for change in human activity at different scales show that the strongest impact of COVID-19 lockdowns were in the transportation sections, and that these impacts varied substantially from country to country. Panel (a) shows the Oxford stringency index (1) for the regions used in this figure. “US (state mean)” is the average of individual states’ indices. “United States” is the index attributed to the US as a whole (not individual states, see SI for discussion). Panel (b) shows the percent change in flights (2-4) for two California airports and three countries (lines) and container moves for three California ports (bars). Panel (c) shows traffic metrics for two California urban areas, and 26 countries (“global”). CalTrans indicates Caltrans PEMS data; Apple indicates Apple driving mobility data. Panel (d) shows electricity consumption in the US by sector, relative to the same month in 2019. The three sectors shown constitute >96% of US power consumption. In (b) and (c), daily metrics are relative to 15 Jan 2020 and presented as 7 day rolling averages and monthly metrics are relative to Jan 2020. Electricity consumption not available after Nov 2020 at time of writing.

Fig. 3. 2020 saw reductions in CO₂, CH₄, and NOₓ emissions. CH₄ and NOₓ are plotted along the left axis, CO₂ on the right. The dashed line for CH₄ after 2017 indicates it is estimated from the average rate of increase. 2020 emissions are represented as a range: the IEA estimated a 10% decrease in CH₄ emissions in 2020 (12), but this is uncertain, as the CH₄ growth rate increased in 2020. Full details are in the SI.

Estimates of the reduction in global NOₓ emissions in the first half of 2020 range from 5% (8) to 13% (7).

The change in global CO₂ emissions was comparable to that of NOₓ emissions, as seen in Fig. 3. Liu et al. report a peak global reduction of approximately 15% (4 Tg C or 15 Mt CO₂) in April, and an annual total of 5.4% (19). In March 2020, Le Quéré et al. projected a slightly larger 7% decrease in CO₂ over the remainder of 2020 (14). The largest decreases occurred in the first half of 2020, as shown in Fig. 4a and were primarily associated with reductions in ground transportation (15). The response of atmospheric CO₂ mixing ratios can be observed near the emissions sources; during the strictest lockdowns, Turner et al. were able to use CO₂ observations from a local ground-based network to estimate a 48% reduction in traffic CO₂ emissions in the San Francisco Bay Area (16). Liu et al. found a 63% (41 ppm) decrease of the typical on-road CO₂ enhancement in Beijing, China (17). Distinguishing these signals in CO₂ at regional scales is more challenging. Buchwitz et al. infer peak decreases in anthropogenic CO₂ emissions from China of 10% from space-based total column CO₂ measurements (18). However, they note that the uncertainty is approximately 100%, and that the expected CO₂ concentration signal is 0.1 to 0.2 ppm, out of a background of over 400 ppm.

Anthropogenic CH₄ emissions are dominated by sources such as landfills, oil and gas production, and agricultural activities. The International Energy Agency (IEA) estimates that CH₄ emissions dropped by 10% in 2020 (Fig. 3), largely due to the decrease in demand for oil and gas. However, it is unclear whether reduced demand during 2020 was the primary driver of emissions. It is likely that decreased maintenance of landfills and oil and gas infrastructure during the COVID-19 pandemic led to new leaks in some areas, which can result in those locations becoming CH₄ “superemitters” (19). In general, the type, maintenance level, and throughput of CH₄ infrastructure can have a large impact of the amount of fugitive emissions. As seen in Fig. 3. Liu et al. report a peak global reduction of approximately 15% (4 Tg C or 15 Mt CO₂) in April, and an annual total of 5.4% (19). In March 2020, Le Quéré et al. projected a slightly larger 7% decrease in CO₂ over the remainder of 2020 (14). The largest decreases occurred in the first half of 2020, as shown in Fig. 4a and were primarily associated with reductions in ground transportation (15). The response of atmospheric CO₂ mixing ratios can be observed near the emissions sources; during the strictest lockdowns, Turner et al. were able to use CO₂ observations from a local ground-based network to estimate a 48% reduction in traffic CO₂ emissions in the San Francisco Bay Area (16). Liu et al. found a 63% (41 ppm) decrease of the typical on-road CO₂ enhancement in Beijing, China (17). Distinguishing these signals in CO₂ at regional scales is more challenging. Buchwitz et al. infer peak decreases in anthropogenic CO₂ emissions from China of 10% from space-based total column CO₂ measurements (18). However, they note that the uncertainty is approximately 100%, and that the expected CO₂ concentration signal is 0.1 to 0.2 ppm, out of a background of over 400 ppm.

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CO₂ and CH₄ atmospheric growth rates

The effect of CO₂ emissions reductions, especially from ground transport, were clearly apparent in urban-scale observations of atmospheric CO₂ mixing ratios (16, 17). This does not, however, transfer to global-scale observations. Figure 4b shows deseasonalized trends in column-average CO₂ mixing ratios (referred to as XCO₂) observed by the Orbiting Carbon Observatory 2 (OCO-2) instrument. Despite the reduction in CO₂ emissions in 2020 (Fig. 4a), there is no clear deflection of the observed XCO₂ below what would be projected based on previous years’ growth rates. We compared the variability in actual atmospheric CO₂ growth rates derived from the OCO-2 data with that computed from fossil fuel emissions (Fig. S8b) and found that the change in atmospheric CO₂ growth caused by the COVID-19 pandemic is smaller than the natural year-to-year variability. This is expected, because the percent change in the CO₂ growth rate, in the absence of feedbacks, will match the percent change in emissions. For a typical growth rate of 2.45 ppm/year since 2016 (Fig. S8b and 22), the 5.4% total reduction in CO₂ emissions calculated by Liu et al. (13) equals a 0.13 ppm/yr decrease in the CO₂ growth rate for 2020—well within the natural variability observed by OCO-2 (Fig. S8) and surface networks (22).

Wildfires are one element of the variability in CO₂ growth rate. The 2019/2020 Australian wildfires emitted 173 Tg C (634 Mt CO₂) between Nov 2019 and Jan 2020, over 6 times more than Australia’s average Nov.-Jan. CO₂ emissions for 2001 through 2018 (23). This drove an early increase in CO₂ in 2020, evident in the deseasonalized southern hemisphere OCO-2 XCO₂ (Fig. 4b, red series) and growth rate derived from the OCO-2 data (Fig. S8b). This wildfire anomaly offset a third of the 518 Tg C (1901 Mt CO₂) reduction in anthropogenic CO₂ (13) and so does not fully explain the offset between emissions and atmospheric mixing ratios for CO₂.

The atmospheric CO₂ growth rate led to a reduction in the rate of oceanic CO₂ uptake. Figure 5 shows the magnitude of ocean carbon fluxes over 8 years as computed from a model ensemble under normal and COVID-like emissions. There is significant variation in the sea-air and CO₂ flux among the model ensemble members. This spread represents the potential interannual variability in CO₂ flux; given that variability, the true change in CO₂ flux in 2020 is uncertain, in part due to corresponding variability in the land carbon sink (Fig. S9). However, the ensemble mean indicates that while on short time scales the land carbon flux is insensitive to the change in emissions (Fig. S9), the ensemble mean ocean uptake was reduced by 70 Tg C/yr in 2020. This would offset 14% of the approximately 520 Tg C/yr (1901 Mt CO₂/yr) reduction in anthropogenic CO₂ emissions in 2020 (13), further dampening the signal from emissions reductions in atmospheric CO₂.

The growth rate of CH₄ was also not slowed by the pandemic. Figure 6a shows trends in column average CH₄ (XCH₄) from two ground based spectrometers in the Total Carbon Column Observing Network (TCCON, 26, 27) located in Park Falls, Wisconsin, US (28) and Lauder, New Zealand (29, 30). The XCH₄ values after 1 March 2020 lie approximately 0.3% above the 2016 to 2019 trend in both hemispheres. Similarly, NOAA reported the single largest increase in CH₄ in its record (31).

Because the lifetime of CH₄ depends on the abundance of...
the hydroxyl radical (OH), the concentration of CH₄ varies with atmospheric pollution levels. In fact, we find compelling evidence that the jump in CH₄ mixing ratios during 2020 is partly due to reductions in NOₓ emissions. In a model incorporating the decreased NOₓ emissions associated with COVID-19 (32), the resulting decrease in global ozone (7) leads to a 2% to 4% decrease in global OH concentrations. As oxidation by OH is the primary loss process for atmospheric CH₄, this acts to increase CH₄ mixing ratios in the atmosphere. Figure 6b compares the trend in XCH₄ measured by TCCON to that predicted by a box model (33). The purple series is the monthly percent difference of TCCON XCH₄ from the linear trends shown in Fig. 6a, and the gray line is the percent difference between a box model run with and without a 3% decrease in OH during 2020. The box model closely matches the extra growth in atmospheric CH₄ during 2020, indicating that the change in OH was an important driver of the observed CH₄ growth. However, this is inconsistent with the 10% decrease estimated by the IEA (12), as our box model assumes constant CH₄ emissions after 2012.

If decreases in anthropogenic NOₓ emissions during 2020 were responsible for the increase in CH₄ lifetime that led to its higher than expected growth rate, what does this imply for the effect of future efforts to reduce NOₓ emissions to improve AQ? To understand this, we need to examine how the 2020 NOₓ decreases affected AQ around the world. In the next section, we will describe the ozone and particulate matter (PM) response to these NOₓ reductions. Afterward, we will explore the implications of this AQ-GHG in the discussion.

**Heterogeneity in air quality response**

Most parts of the world saw significant decreases in NOₓ emissions during the pandemic, but the magnitude and timing of these emissions changes varied with location. Figures 7a-c compare timeseries of NO₂ column densities measured by TROPOMI for three cities. Following the beginning of lockdown measures (indicated by the dotted lines), the 2020 NO₂ column densities are clearly less than in 2019. However, in Los Angeles, the drop in NO₂ occurred very rapidly when lockdowns were enacted in early March, but by May there was little difference between 2019 and 2020. In Lima, on the other hand, the difference between 2019 and 2020 grew from March through May. In Shanghai, we see a very large drop in NO₂ associated with the early lockdown in January and a smaller drop during the second lockdown in late February.

These changes in NOₓ emissions drove changes in secondary pollutants, such as ozone and PM. However, the ozone and PM responses depended on the local chemical regime and meteorology, as well as the magnitude and timing of the NOₓ emissions reductions. In this section, we describe the factors controlling the ozone response first, followed by PM.

**Ozone.** Ozone is a secondary pollutant produced in the atmosphere from the reaction of NOₓ and OH with volatile organic compounds (VOCs). The response of ozone concentrations to changes in NOₓ emissions is characterized by the ozone production efficiency (OPE), which is the ratio of the change in ozone for a given change in NOₓ. More detail is given in the SI.

Two patterns in the OPEs demonstrate the significant spatial and temporal variability in the relationship between NOₓ emissions and ozone concentrations. First, in Fig. 7f, the OPE in the Northern hemisphere increases between February and June. This is mostly due to increasing sunlight driving key photolysis reactions more rapidly. Thus, the timing of NOₓ emissions changes plays a significant role in the magnitude of the ozone response in the mid- and high-latitudes, with a smaller ozone response to a given NOₓ change during spring than during summer. Second, in Fig. 7d, tropical and subtropical cities have the largest, most positive OPEs. Furthermore, there is little change in OPE with season for these cities (Fig. 7e) due to the relatively small changes in insolation at low latitudes. Figure 7d indicates that most of the northern mid-latitude cities have small, positive OPEs. Two cities, however, have slightly negative OPEs (Beijing -0.10, Karachi -0.06); a negative OPE indicates that ozone increased when NOₓ emissions decreased. Other studies have, in fact, identified large ozone increases in China (34) associated with the decreased NOₓ emissions during the pandemic. Additional increases in ozone were observed in Europe (35), with smaller but still positive changes in ozone in the United Kingdom (36).

We use a steady-state model (Fig. S10) to interpret the patterns in Fig. 7. From the steady-state model, we know OPE is small at both low and high NOₓ concentrations, but large at...
intermediate NO\textsubscript{x} concentrations. Overall OPE also increases with VOC reactivity (VOC\textsubscript{R}), the total rate of reaction of all VOCs with OH in a given parcel of air) for NO\textsubscript{x} concentrations greater than \( \sim 0.1 \) ppb. Thus, in Fig. 7, areas with negative OPE are in the high-NO\textsubscript{x} part of the OPE curve; sustained efforts to reduce NO\textsubscript{x} emissions will bring them closer to the maximum-OPE tipping point, after which NO\textsubscript{x} reductions should lead to ozone reductions. Cities in the tropics and subtropics have large, positive OPE values. This is partly due to plentiful sunlight to drive photochemistry, but these regions also have large VOC\textsubscript{R} values due to the abundance of biogenic VOCs (37). The steep dependence of OPE on NO\textsubscript{x} follows because NO\textsubscript{x} is the limiting reactant in ozone production in these high-VOC\textsubscript{R} conditions. Thus, these cities should see large ozone reductions from NO\textsubscript{x} reductions. However, of the equatorial cities shown in Figure 6, only those located in South Asia had large enough reductions in NO\textsubscript{x} emissions during the COVID-19 pandemic to produce substantial reductions in surface ozone (3-5 ppb) (7).

We also see this heterogeneity in ozone response to NO\textsubscript{x} emissions reductions at the intraurban scale. Measurements of daily maximum NO\textsubscript{x} and ozone at monitoring sites throughout the Los Angeles Basin show consistent reductions in NO\textsubscript{x} throughout the basin in March and April of 2020, but smaller reductions in ozone in the central northern part of the basin than elsewhere (Figs. S1, S2). This is consistent with the near-0 OPE for Los Angeles in Fig. 7d, i.e., for a city on the verge of reducing NO\textsubscript{x} emissions to the point where NO\textsubscript{x} is the limiting factor in ozone production. While the overall basin chemistry is at this tipping point, local differences in emissions as well as transport of pollutants within the basin can lead to these small scale differences in ozone response (38).

However, the behavior of ozone in the Los Angeles Basin also illustrates that NO\textsubscript{x} controls may become less effective in a warmer climate. Figure 8 shows time series of daily maximum NO\textsubscript{x} and ozone (top and middle panels). NO\textsubscript{x} and ozone concentrations are clearly lower in March and April 2020 compared to the 2015 to 2019 average, in part due to the reduction in NO\textsubscript{x} emissions at the beginning of the lockdown. However, these two months were significantly cooler than the 2015 to 2019 average as well. When temperatures rose above average during an unusual heat wave in late April and May of 2020, ozone daily maxima rose above the range seen in 2015 to 2019, despite the fact that NO\textsubscript{x} remained similar to 2015 to 2019 concentrations. An increase in ozone during April and May was also seen in a previous study (39). The response of ozone per degree increase in temperature is shown in Fig. S3. Typical values for the O\textsubscript{3} season (May-Sep) in 2020 throughout the basin were 1.8 to 5.8 ppb K\textsuperscript{-1}. This is higher than a previous prediction of about 1 ppb K\textsuperscript{-1} in the basin (40), suggesting the ozone climate penalty may be stronger than expected; however, analysis is ongoing.

**Particulate matter.** Achieving long-term reductions in PM (especially PM 2.5, particles with a diameter \(< 2.5 \mu m\)) concentrations is a matter of great importance due to the large health impacts of PM compared to ozone (41). Our interest here is to use observations from the pandemic period to better understand some of the factors controlling atmospheric PM concentrations, rather than focusing on the question of whether
PM exposure increases the chance of death from COVID-19. The factors controlling PM concentration are more complicated than those for ozone. PM arises from primary emissions and natural sources, as well as secondary chemistry in the atmosphere. One such secondary pathway is the formation of nitrate PM from the reaction of higher oxides of nitrogen (such as HNO₃) with ammonia (42). Nitrate PM formation via this pathway may be limited by either available NOₓ or ammonia.

Model simulations (Fig. S4) demonstrate the effect that NOₓ emissions reductions had on nitrate PM formation in Los Angeles. Under COVID-19 emissions, the nitrate PM concentrations decreased by approximately 60% in April 2020. At the same time, the model reported a shift towards NOₓ-limited (rather than ammonia-limited) chemistry. This implies that the NOₓ emissions decreases in April, when the shift in the chemical regime shows the largest change, were more efficient at reducing nitrate than the reductions in other months. Compared to the measured total PM reductions shown in the bottom panel of Fig. 8, our results suggest that NOₓ emissions reductions account for about 10% of the total PM reduction in the Los Angeles Basin during the COVID-19 lockdowns. This agrees with other recent work (43) which indicate that traffic NOₓ emissions contribute less than 10% of secondary PM production throughout North America, Europe, and East Asia.

The relative availability of NOₓ and ammonia elsewhere in the US plays an important role in whether NOₓ emissions reductions lead to reduced nitrate PM. Simulations of nitrate chemistry over the continental US show that Los Angeles is somewhat unique as an urban area that experienced a significant shift to NOₓ-limited nitrate chemistry. Other urban areas in the northeast, southeast, and northwest largely remained ammonia-limited (Figs. S5–S7). This could explain, at least in part, the scattered response of PM to NOₓ emissions reductions across US cities seen in other studies (44). It also implies that continuing the long-running trajectory of NOₓ emissions reductions in Los Angeles in order to reach the tipping point where ozone becomes NOₓ limited will also benefit AQ via reduced production of nitrate PM.

However, Los Angeles also represents a cautionary tale about attributing AQ changes to the COVID-19 pandemic without accounting for other confounding factors. Weather and wildfires also played a large role in determining the PM concentrations in Los Angeles during 2020. When the lockdowns were first instituted in March, news outlets and social media attributed the clear air in the Los Angeles Basin to the lack of traffic. However, as seen in Fig. 8, the lower PM concentrations in March and April 2020 than 2015 to 2019 (Fig. 8, bottom) coincide with anomalously cool weather, which was accompanied by higher than average precipitation (Fig. S1 in (38)). Precipitation removes PM from the atmosphere through wet deposition (45, 46), and was at least partially responsible for the clean air during this period. The extreme spike in PM concentrations seen in September 2020, on the other hand, coincides with a time period when major wildfires were burning in close proximity to Los Angeles. Like the April-May heatwave, this event also points to the fact that climate change can erase progress in AQ improvement through emissions reductions.

**Discussion**

The changes in atmospheric composition throughout 2020 unequivocally demonstrate that AQ and GHGs cannot be treated as separate problems, despite the disparate time scales of AQ and GHG responses to changes in human activity. AQ is most dependent on local changes in emissions because of the shorter atmospheric lifetime and rapid chemistry of AQ-relevant pollutants. In contrast, the global total GHG emissions matter more than local emissions, as it is the overall GHG atmospheric growth rate that drives climate change. As discussed above, improvements in AQ made by reducing pollutant emissions locally can be offset by changes in meteorology or non-anthropogenic (e.g., biogenic or wildfire) emissions driven by climate change. Likewise, changes in AQ can affect climate change, as decreases in AQ-relevant emissions could lead to increased lifetimes for shorter-lived GHGs (such as CH₄), increasing their global warming potential.

Reductions in NOₓ emissions during the pandemic did show the potential benefits cities can gain by promoting systemic change to accomplish these same reductions. For most countries, the pandemic-induced emissions reductions can be seen as going back in time to a period when NOₓ emissions were lower. In the US, Europe, and China, where NOₓ emissions have been trending downward, these reductions were more
COVID-19 Equivalent NOx Emissions Year by Country

Fig. 9. The emissions reductions during the pandemic are, in a sense, like moving forward or back in time. Countries are colored by the year to which their 2020 NOx emissions are equivalent, projected forward in time where emissions have been decreasing and backward elsewhere. Details of emissions estimates given in the SI.

...emissions hold constant. Most striking is how much more quickly China could reach pandemic-like emissions levels than the US or Europe. Though all three regions’ emissions reductions had similar peak magnitudes (18% to 20%), Europe and especially the US are further along their respective NOx reduction pathways than China. This, combined with China’s higher pre-pandemic emissions levels, means that China can make progress quickly if they are able to maintain the aggressive pace of emissions reductions they have set over the past decade (32).

Many cities in the US and Europe are close to reaching a point at which NOx emissions will be a very effective control on ozone concentrations. In Fig. 7d, cities with an OPE near 0 are likely at the tipping point between VOC-limited and NOx-limited chemistry. Further NOx reductions should move them firmly into NOx-limited chemistry, where NOx is the primary control on ozone formation. While sustaining these emissions reductions may be challenging due to the decreasing contribution of on-road gasoline emissions (47) and the impact of emissions reductions being offset in part by increases in chemical lifetime (48), the rewards in doing so are likely substantial. In addition, since NOx and CO2 are co-emitted by combustion processes, regulations such as those that encourage a transition to electric vehicles will also benefit climate. In fact, recent work has shown that the cost savings associated with reduced health impacts from air pollution will outweigh the cost of transition to a clean carbon economy and that the increased radiative forcing from longer-lived CH4 and ozone is balanced by the decrease in forcing from smaller CO2 mixing ratios (49). On the other hand, measures such as NOx removal from coal-fired power plants will benefit AQ but not climate; as discussed below, this will eventually limit their effectiveness for improving AQ.

The same strategies to improve AQ will not be equally effective in all locations. On one hand, the tropical and subtropical cities with large, positive OPE values in Fig. 7d can immediately realize substantial ozone reductions through reductions in NOx emissions. On the other hand, cities such as Beijing and Karachi with negative OPEs, or locations such as the United Kingdom where in situ studies found a negative correlation between NOx emissions and ozone concentrations (36) would do better to reduce volatile organic compound (VOC) reactivity simultaneously with NOx emissions. Such an approach would allow them to avoid the chemical regimes with the largest OPEs (50) (Fig. S10a). Similarly, while chemical formation of ammonium nitrate PM in Los Angeles became NOx-limited during the pandemic, most other cities in the US remain ammonia-limited and would see stronger reductions in PM by controlling primary emissions, organic precursors, or other key species.

Unfortunately, 2020 has also shown that improvements in AQ are likely to be offset by climate feedbacks. Such effects were most apparent in Los Angeles, where warmer than average May temperatures led to ozone concentrations above the 2015 to 2019 average, greater than average precipitation in March and April likely contributed to the reduction in PM, and severe wildfires from late August through September caused PM concentrations four times that of the 2015 to 2019 average. Changing climate will affect each of these variables, leading to warmer temperatures, more wildfires (51), and potentially more intense but less frequent precipitation (52), giving PM more time to accumulate between wet deposition events.

Changes in AQ-relevant emissions, particularly NOx emissions, have potential to feed back into climate change as well. As we showed in Fig. 6, there is compelling evidence that reductions in OH stemming from reduced anthropogenic NOx emissions drove a ∼0.3% jump in CH4 during 2020. While tropical cities have the greatest potential for decreasing ozone by reducing NOx emissions (Fig. 7d), they also have an outsized impact on atmospheric CH4 lifetime, as the largest share of CH4 oxidation occurs in the tropics (33). Since only tropical cities in South Asia had substantial changes in NOx emissions during 2020 (7), 2020 represents a minimum benchmark for the effect of NOx reductions on the CH4 growth rate. It is therefore essential to invest strategies to reduce fugitive CH4 emissions (such as updated CH4 storage and transportation infrastructure to prevent and limit leaks, landfill CH4 capture, and confined animal feed operation CH4 mitigation) ahead of decreases in tropical NOx emissions.

In terms of climate, despite a reduction in global emissions equivalent to going back in time nine years (to 2011-equivalent CO2 emissions), any change to the global CO2 growth rate was smaller than typical interannual variability. As mentioned earlier and discussed in more detail below, this is partly due to the offsetting reduction in ocean carbon uptake (Fig. 5), but also arises because the sharp decreases in CO2 emissions during the first half of 2020 were not sustained. By the second half of 2020, emissions due to power generation, industry, and residential consumption had nearly returned to 2019 levels (13). If we assume that these emissions levels represent a balance between reduced activity to limit the spread of COVID-19 and sufficient activity to maintain a minimum economic productivity, this suggests that reducing activity in these sectors is not practical. Reducing these sectors’ emissions permanently will require their transition to low carbon emitting technologies.

One interesting aspect of the GHG emissions reductions during the pandemic was that they provided a chance to study
the feedback in ocean carbon uptake. The model simulations using COVID-like CO$_2$ emissions shown in Fig. 5 indicate that the sea-air carbon flux adjusts rapidly in response to changes in anthropogenic emissions. That model ensemble mean indicates a response time of about one year. Though this basic response - a decline of the ocean carbon sink in response to mitigation - is accounted for the RCP scenarios (53), much uncertainty remains as to the accuracy of these ocean sink predictions. This uncertainty is due both to the forced response of the ocean and to interannual variability Lovenduski et al. found that, for a change in ocean carbon uptake to be observable with our current network of ocean buoy measurements, it would need to be four times larger than the COVID-19 emissions reductions (25). This will be a challenge as we work to quantify the effect of future permanent CO$_2$ emissions reductions on atmospheric CO$_2$ mixing ratios.

The pandemic does offer insight into how the atmospheric GHG growth rates could be curtailed: systemic changes are required to enable sustained reductions in emissions. The efficacy of sustained reductions (without systemic changes to the energy sector) can be seen in the contrast between CO$_2$ emissions from ground transport and international shipping and aviation (“international bunkers”) reported by Liu et al. (13) The peak reduction in international bunkers’ emissions was only approximately 1/3rd that of the reduction in emissions from ground transport, by mass. However, while ground transport recovered fairly quickly, the international bunkers’ emissions remained at about half of 2019 levels throughout the second half of 2020. As a result, the cumulative reduction in 2020 emission due to international bunkers was 75% that of the reduction due to traffic, despite the comparatively small magnitude of the daily emissions from international bunkers.

Sustained reduction in other sectors will require investment in renewable energy and new technologies to support current levels of productivity with lower carbon emissions. That is, to reduce the carbon intensity of our economy. Such investment is essential, as several studies (54, 55) have documented the harm to employment, family connections, and other critical human connections from the reduction in personal mobility due to the pandemic. Liu et al. (13) note that Spain’s 2020 emissions due to power generation were almost 25% lower than in 2019 due to investment in renewable energy. A post-COVID economic recovery represents an opportunity to invest in carbon-reducing technologies (56), as long as the need to balance short-term job creation with long-term retraining is accounted for (57). If this investment was able to continue the trend of a 5.4% decrease in global CO$_2$ emissions per year, we would reach “preindustrial” (circa 1850) emissions levels in approximately 18 years.

**Strengths and weaknesses of current observing systems**

Understanding how the COVID-19 pandemic has altered AQ and the carbon cycle has relied heavily on the multifaceted observing system built over the past two decades, including satellites, dense ground-based observing networks, Earth system and chemical transport models, and techniques to assimilate observations into these models. Novel data on human activity (particularly internet-of-things mobility data, crowdsourced air traffic data, and even news reports) have also played a vital role in both understanding how human behavior changed during the pandemic and quantifying the effect of that change on anthropogenic emissions.

Nevertheless, there remain important gaps in our observing network. First, space-based detection of VOCs remains a challenging problem, yet quantitative measurements of key biogenic (e.g. isoprene, terpenes) and anthropogenic (e.g. ethene, propene) contributors to VOC OH reactivity are needed to identify the dominant chemistry governing AQ around the globe. Second, as we saw in the LA Basin case study, disentangling primary PM emission, secondary PM formation, and meteorological drivers of PM concentration is crucial to understand which processes control PM exposure. Given the serious health impacts of PM exposure, work towards an integrated surface and space-based system that can differentiate these processes is needed to elucidate the optimum approaches to reducing PM exposure.

In regards to climate-relevant observations, spatiotemporally broader and denser space-based GHG observations would provide a highly valuable empirical constraint on changes to anthropogenic and biogenic carbon fluxes. A satellite instrument that provided comparable observations to the BEACO2N network in the San Francisco Bay area (∼ 2 km resolution, strong sensitivity to the near-surface atmosphere, urban-scale coverage) could apply similar inversion techniques as Turner et al. (16) to infer key sectors’ emissions in cities around the world. It is also clear that our current network of near-real time ocean carbon uptake measurements are not sufficient to disentangle internal variability in the air-sea carbon flux from changes driven by reductions in anthropogenic emissions (25). Expanding this network or developing new methods to constrain the air-sea carbon flux from space will be necessary to quantify the impact of anthropogenic emissions reductions on atmospheric CO$_2$ mixing ratios.

**Conclusions**

The COVID-19 pandemic and associated changes in human behavior represent an unprecedented rapid change in anthropogenic emissions to the atmosphere. Due to the large differences in relevant atmospheric lifetimes for constituents central to AQ and climate, clear changes in local AQ but not global GHG trajectories were observed. Changes in AQ were very spatially heterogeneous, demonstrating that the same strategies to improve AQ do not apply equally well to all regions. Additionally, changes in AQ in the Los Angeles Basin correlated with temperature, precipitation, and severe wildfires, indicating that shifts in these quantities associated with climate change will at least partially offset gains in AQ made from past and future reductions in anthropogenic emissions. Despite large disruptions in transportation sectors, the global-scale change in the CO$_2$ growth rate was less than interannual variability. This is due to a combination of reduced ocean uptake of CO$_2$, a recovery of CO$_2$ emissions in the second half of 2020, and large interannual variability in land carbon fluxes. That recovery indicates that expecting changes to individual behavior to be sufficient to halt the increase of GHGs in the atmosphere is unrealistic. Instead, incentives to deploy new methods to systematically and sustainably reduce carbon intensity are needed. Given the bidirectional feedback between climate and AQ, it is clear that climate and AQ can no longer be considered separate problems; prompt action to reduce anthropogenic carbon emissions is essential not only to
affect direct climate impacts, but to avoid giving up decades of hard-won progress in improving urban AQ.

Materials and Methods

Full methods are available in the SI. Analysis of LA Basin AQ used data from CA Air Resources Board monitors, filtered for complete data records in the 2015 to 2020 period. 1 h daily maximum (DM) NO₂ and temperature, 8 h DM O₃, and 24 h average PM were calculated from this data. OPE was derived from model simulations using multicoinstuent assimilation of multiple satellite measurements in the MIROC-CHASER model (32). OPE was calculated by comparing modeled O₃ production and NO₂ emission difference between baseline (2010 to 2019) and reduced 2020 emissions. Separate PM₂.₅ simulations used GEOS-Chem v9-02 with NO₂ emissions consistent with the OPE simulations: baseline NO₂ emissions used HTAP v2 scaled to 2017 using satellite-derived emissions reduction ratios and COVID NO₂ emissions were scaled down by the same factor as in the OPE simulations. The TROPOMI timeseries analysis first regressed native TROPOMI pixels to a 0.01° × 0.01° grid and filtered to primarily remove cloud and snow/ice contaminated scenes. The timeseries show the 75th percentile of 15-day rolling average NO₂ columns in a 1° × 1° box around each city.

Global CO₂ emissions estimates were derived from an array of near-real time data on power generation, industry, transport, and fuel consumption. XCO₂ growth rates were derived from OCO-2 v10 ocean glint data and XCH₄ growth rates from TCCON GGG2014 data. The data shown are 15-day running averages deseasonalized by fitting a four-harmonic curve. Expected CH₄ trends we computed from a two-box model (representing the two hemispheres) using prescribed OH concentrations and constant CH₄ emissions after 2012. TCCON data can be obtained from the TCCON Data Archive hosted by CaltechDATA (https://tcconda.org/). The authors thank the TCCON science team for their effort in providing this data.

Publicly available datasets are listed along with data generated from this study and stored in public facing repositories in the SI, table S1. Emissions data for Figs. 3 and 9 are given in Table S2. Data for the OPE values in Fig. 7 is given in Table S4. Emissions and OPE data also included as Excel SI files.

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Supplementary Information for

Societal shifts due to COVID-19 reveal large-scale complexities and feedbacks between atmospheric chemistry and climate change


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This PDF file includes:
Supplementary text
Figs. S1 to S11
Tables S1 to S4
References for SI reference citations
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Fig. S1. Change in 1 hr daily maximum (DM) NO$_2$ in 2020 relative to the average of 2015 to 2019 at the California Air Resources Board sites throughout the South Coast Air Basin.

Fig. S2. Change in 8 hr daily maximum (DM) O$_3$ in 2020 relative to the average of 2015 to 2019 at the California Air Resources Board sites throughout the South Coast Air Basin.
Fig. S3. Average derivatives of $O_3$ response vs. temperature between May and September at California Air Resources Board sites throughout the South Coast Air Basin for years 2015–2020. Each group of bars is one site, and are ordered by longitude (west to east).

Fig. S4. Simulated inorganic nitrate aerosol sensitivity at downtown LA for two model runs during March to May 2020. Dashed lines represent the run with lockdown-induced emissions reductions (COVID-19), solid lines represent the business as usual (BAU) run. NO$_x$ emissions are shown in black, nitrate aerosol concentration in blue, and the gas ratio in red. A gas ratio $< 1$ indicates NH$_3$-limited (compared to NO$_x$-limited chemistry). See the SI for more information.
Fig. S5. Average change in gas ratios for March 2020 between a model simulation using business as usual (BAU) NOx emissions and one using emissions based on NO2 observations for March 2020 (COVID-19). The gas ratio is described in Eq. (3); a value < 1 indicates NH3 limited nitrate aerosol formation; a value > 1 indicates NOx limited aerosol formation.

Fig. S6. Same as Fig. S7, but for April 2020.
Fig. S7. Same as Fig. S7, but for May 2020.
Fig. S8. Trends in column average CO₂ and CH₄. (a) Trends in CO₂ from the Orbiting Carbon Observatory 2 (OCO-2) for the northern and southern hemispheres. The pale blue and red markers are daily values, calculated as described in the text. The vibrant blue and red markers represent deseasonalized values computed from the daily values by fitting a fixed seasonal cycle described by a four-term harmonic equation (1, 2). The solid line is a robust linear fit to the 2016 through 2019 data. (b) Annual growth rate of CO₂ computed from OCO-2 data in the northern and southern hemispheres, as well as derived from fossil fuel emissions trends. See text for details. (c) As (a), but for CO₂ from two Total Carbon Column Observing Network (TCCON) stations: Park Falls, WI, USA in the northern hemisphere and Lauder, New Zealand in the southern hemisphere. (d) As (b), but derived from TCCON CO₂. (e) As (c), but for CH₄. (f) As (d), but for CH₄. In all panels the vertical gray dashed line marks 1 March 2020.
**Fig. S9.** Annual mean, globally integrated terrestrial net ecosystem production (NEP, positive into biosphere, excludes land use change) predicted from the CanESM5-COVID ensemble (3). As in the main paper, black/gray lines derive from simulations forced with SSP2-RCP4.5 CO$_2$ emissions, while red/pink lines derive from simulations forced with a 25% peak CO$_2$ emissions reduction in 2020. See (3) for more details. Thick lines are ensemble averages, and thin lines are individual ensemble members, each with different phasing of internal variability.

**Fig. S10.** Theoretical ozone production efficiency as a function of NO$_x$ concentration and one other variable, computed in a steady-state model. In all panels, NO$_x$ concentration is given on the $x$-axis, the second independent variable on the $y$-axis, and the color represents the ozone production efficiency. In panel (a), the $y$-axis is total VOC reactivity, VOC$_R$; in panel (b), it is total HO$_x$ production, P(HO$_x$); in panel (c), it is the branching ratio ($\alpha$) for the RO$_2$ + NO reaction. Note that the $y$-axis in panel (b) is multiplied by $10^7$ and the color scale for panel (c) has a higher maximum value than the other panels and is logarithmic, rather than linear. The default values for VOC$_R$, P(HO$_x$), and $\alpha$ when not the second dependent variables are $5.0 \text{ s}^{-1}$, $6.25 \times 10^6 \text{ molec. cm}^{-3} \text{ s}^{-1}$, and 0.04, respectively.
Supporting Information Text

Methods

Public data. All public datasets used in this study are shown in Table S1.


The CAADA Python package (26) was used to preprocess the PeMS vehicle counts and Strohmeier et al. (5) flight data, as well as download Port of LA and Port of Oakland container moves. For the purposes of Fig. 2, “Bay Area” is defined as Alameda, Contra Costa, Marin, San Mateo, San Francisco, Santa Clara, and Santa Cruz counties, while “LA” is defined as Los Angeles, Orange, Riverside, San Bernardino, Santa Barbara, and Ventura counties. For flight data, shipping data, and traffic data, daily values were normalized such that 15 Jan 2020 is 100% and monthly values were normalized such that Jan 2020 was 100%. For electricity use data, each month’s value is the 2020 use as a percentage of 2019 use in the same month.

**Oxford stringency index: US vs. US state mean.** The Oxford Stringency Index (27) includes stringency metrics labeled as US without a subregional code along with metrics for individual states. In Fig. 2, “United States” indicates that the US values without a subregional code are plotted, while “US (state mean)” indicates that the average of all the individual states’ stringency indices is plotted. The Oxford index subnational interpretation guide (https://github.com/OxCGRT/covid-policy-tracker/blob/master/documentation/subnational_interpretation.md, last accessed 13 May 2021) indicates that their primary dataset summarizes the totality of policies in the specified territory.

While we include both the combined US and state mean metric to illustrate the general stringency of lockdown measures in the US, we do not ascribe specific meaning to the difference between them.

**Equivalent Emissions Year Calculations.** For the CO2 emissions in Fig. 3, we used 2005-2018 fossil fuel emissions from the Global Carbon Budget 2019 (28). For 2019, we assumed a +0.1% increase from 2018 based on Supplementary Data in Le Quere et al (29). For 2020 we used a 7% decrease from the 2019 value with a ±1% uncertainty, based on Le Quere et al (29) and Liu et al (30). The 2020 emissions are 9.29 (± 0.10) GtC/yr; this corresponds to somewhere between 2010 (9.05 GtC/yr) and 2012 (9.50 GtC/yr). For CH4, we use the anthropogenic emissions based on the EDGARv4.3.2 and GfED4.1s emissions inventories as published in the Global Methane Budget 2000-2017 (31). To estimate the emissions trajectory beyond 2017, we assumed that the rate of increase for 2018 and 2019 was equal to the average rate for 2005 to 2017, then used the estimated 10% reduction in 2020 from (32). For the global NO2 emissions trajectory in Fig. 3 we used 2005-2020 emissions from the assimilation system described in the subsection “Global ozone production efficiency calculation” below.

For Fig 9, we again used the NO2 emissions from the assimilation system. For countries whose emissions have been monotonically increasing since 2005, we calculate the prior year with the same emissions as 2020. For countries whose emissions decreased over all or part of the 2005-2019 period, we use the 2015-2019 rate of decline to project emissions into the future.

**Global CO2 emissions estimates.** We calculated the daily global fossil CO2 emissions in 2020 (updated to December 31st), as well as the daily sectoral emissions from power sector, industry sector, transport sector (including ground transport, aviation and shipping), and residential sector respectively. The estimates are based on a set of near real time dataset including hourly to daily electrical power generation data from national electricity operation systems of 31 countries, real-time mobility data (TomTom city congestion index data of 416 cities worldwide and FlightRadar24 individual flight location data), monthly industrial production data (calculated separately by cement production, steel production, chemical production and other industrial production of 27 industries) or indices (primarily Industrial Production Index) from national statistics of 62 countries/regions, and monthly fuel consumption data corrected for the daily population-weighted air temperature in 206 countries.

**CO2 and CH4 trends and CH4 box model.** CO2 and CH4 trends were computed from version 10 column average CO2 (termed XCO2) measurements made by the Orbiting Carbon Observatory 2 (OCO-2) satellite instrument and ground based CO2 and CH4 column measurements from two Total Carbon Column Observing Network (TCCON) sites: one in Park Falls, WI, USA (45.945° N, 90.273° W) and Lauder, New Zealand (45.038° S, 169.684° E). OCO-2 data was subset to quality flag = 0 data collected in the ocean glint mode and all data averaged daily between 20° N and 55° N for the northern hemisphere and 55° S and 20° S for the southern hemisphere. TCCON data was limited to data with flag = 0; publicly available data is already filtered in this manner.

To compute the trends, 15 day running averages of the daily data were computed and deseasonalized using the method in Liu et al. (1) which follows Graven et al. (2). A robust linear fit was applied to the 2016 through 2019 data. 2020 was excluded so as to test how the 2020 trend compared to the previous four years.

Growth rates were computed from the deseasonalized data by taking the differences in time of three month averages of the OCO-2 or TCCON deseasonalized data, multiplied by four to convert from three-monthly to annual growth rates. The growth
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Table S1. Public data sources used in this paper. The “Used for” column gives the part of the analysis in which that data was used.
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Table S2. Emissions used in Figs. 3 and 9. A dash indicates that emissions data were not available for that year.
rate for fossil fuel emissions was computed by using three month total of anthropogenic CO\(_2\) emissions from the Open-source Data Inventory for Anthropogenic CO\(_2\) (ODIAC) for 2016 through 2019 and carbonmonitor.org for 2020. The three month total emitted CO\(_2\) mass was converted to an atmospheric mixing ratio by:

\[
R_{FF} = 4 \cdot E_{CO2,3mo} \cdot \frac{2.14\text{ ppm}}{\text{Gt C}} \cdot f
\]  

where \(E_{CO2,3mo}\) is the three month total CO\(_2\) emissions and \(f\) is the average airborne fraction computed from all of the OCO-2 data; each three-monthly airborne fraction \((f)\) is computed as:

\[
f = \frac{R_{OCO-2,3mo}}{E_{CO2,3mo} \cdot 2.14\text{ ppm/Gt C}}
\]

where \(R_{OCO-2,3mo}\) is the three-monthly growth rate computed from the OCO-2 data.

The TCCON CH\(_4\) series shown in Fig. 6b are computed from the time series and trends in Fig. 6a. First, the percent difference of the northern and southern hemisphere data against their respective trends is computed. Then, monthly averages of these two percent differences are calculated. Finally, the two monthly time series are averaged together.

The box model trend shown in Fig. 6b was calculated using the box model described in (33) and (34), available at https://github.com/alexjturner/BoxModel_PNAS_20161223. Briefly, this model treats the change in concentration of CH\(_4\) in each hemisphere as the sum of changes due to emissions, oxidation by OH, and interhemispheric transport. OH concentrations can either be directly prescribed or have a prescribed source with concentrations varying alongside CH\(_4\) and CO. The results in Fig. 6b use prescribed OH concentrations, but the behavior is similar if the OH source is prescribed. For simplicity, CH\(_4\) emissions followed the “stabilized” scenario described in (34). The percent difference in CH\(_4\) shown in Fig. 6b is the difference between a model run with a 3% reduction in OH during 2020 and one without.

We do note that, in the box model, the renewed CH\(_4\) growth after 2008 occurs earlier than indicated by in situ measurement. This is due to the timing of CH\(_4\) emissions growth in the EDGAR inventory. However, this does not affect our conclusions as (a) we use the difference of two model runs with the same CH\(_4\) emissions trends and (b) we focus on the behavior in 2020.

**TROPOMI NO\(_2\) timeseries.** For our analysis we re-grid the operational TROPOMI tropospheric vertical column NO\(_2\), with native pixels of approximately 3.5 x 7 km\(^2\) for 2019 and 3.5 x 5.5 km\(^2\) for 2020, to a newly defined 0.01° x 0.01° grid (approximately 1 x 1 km\(^2\)) centered over each of the three cities: Los Angeles, Lima, and Shanghai. Before re-gridding, the data are filtered so as to use only the highest quality measurements (quality assurance flag (QA_flag) > 0.75). By restricting to this QA value, we are removing mostly cloudy scenes (cloud radiance fraction > 0.5) and observations over snow-ice. Once the re-gridding has been completed, the data is binned temporally during a 15-day rolling timeframe and spatially over the metropolitan area, which we loosely define as a 1º x 1º box over the city center. The rolling 75th percentile of the binned data during the first five months of 2019 and 2020 are shown in top row of Figure 7. There is some evidence that the current TROPOMI operational NO\(_2\) product may have a low bias of 20 to 40% in polluted areas; much of this bias may be attributed to the air mass factor (35–37). We limit our analysis to relative trends, which reduces this uncertainty.

**LA Basin AQ analysis.** The hourly ambient temperature and concentrations of PM\(_2.5\), NO\(_2\), and O\(_3\) in the South Coast Air Basin for the period of 1 Jan 2015 to 30 Sept 2020 were downloaded from the California Air Resources Board Air Quality Data Query Tool (https://www.arb.ca.gov/aqmis2/audsselect.php). It should be noted that the 2020 data are preliminary, unvalidated, and subject to change. The following steps were taken for data analysis:

1. Only the monitoring sites that had complete data between 2015 and 2020 were considered in this analysis. Near-road monitoring sites were not included in the analysis. Figure S11 and Table S3 show the location of the monitoring sites considered in this analysis and the parameters measured at each site, respectively.

2. For every date and site, the 1hr daily maximum (DM) temperature, 24hr average PM\(_2.5\), 1hr DM NO\(_2\), and 8hr average DM O\(_3\) were calculated.

3. For every date, the average of the above-mentioned parameters was calculated across all monitoring sites. 7-day moving averages were then calculated and presented by day of year in Figure 8 for 2020 and the average (± range) of [2015-2019]. The background colors in Figure 8 illustrate the difference between the 7-day moving average temperature in 2020 and the average (±1σ) temperature in [2015-2019] by day of year.

4. Using the data in step 2, the percent change in monthly average concentrations of 1hr DM NO\(_2\) and 8hr DM O\(_3\) between 2020 and the average of [2015-2019] was calculated by month and site as shown in Figures S1 and S2.

**Global ozone production efficiency calculation.** We evaluated the seasonal and regional changes in the global tropospheric ozone response to COVID-19 NO\(_x\) emissions using a state-of-the-art chemical data assimilation system. Anthropogenic NO\(_x\) emission reductions linked to the COVID-19 pandemic were estimated as the difference between 2020 emissions and climatological (baseline) emissions for 2010-2019 estimated from our decadal chemical reanalysis constrained by multiple satellite measurements. The assimilation system uses the MIROC-CHASER global chemical transport model and an ensemble Kalman filter technique (38). This approach allows us to capture temporal and spatial variations in transport and chemical
Fig. S11. Location of South Coast Air Basin monitoring sites included in this analysis.
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Table S3. Parameters used from each South Coast Air Basin monitoring site.
204 reactions in the emission and concentration estimates. The results for 2020 were used previously to evaluate the air quality
response to Chinese COVID-19 lockdown (39), and show reasonable agreements with the observed concentrations from in-situ,
ozonesonde, and satellite ozone measurements globally for 2005–2018 (39) as well as for 2020 (40).

In order to evaluate seasonal and regional differences in the ozone response, the ozone production efficiency (OPE) was
estimated based on model sensitivity calculations using the 2020 and baseline emissions for February-July 2020. The OPE was
calculated using the simulated global tropospheric ozone burden changes corresponding to changing NOx emissions (i.e., the
COVID-19 emission anomaly); the analysis was performed separately for each of the selected megacities. The model simulations
were conducted from the beginning to the end of each month for the time period February to June, 2020, using the same initial
conditions. The simulated tropospheric ozone burden averaged over the last 5 days of each month was compared between the
simulations using the 2020 and baseline emissions. The analysis thus provides information on monthly changes in the ozone
response (Tg) to reduced NOx emissions (Tg per year) for each megacity separately. These data are presented in Table S4.

PM2.5 simulations. We used the GEOS-Chem (v9-02) model with a bi-directional NH3 flux scheme (41) at the nested resolution
of 0.3125° × 0.25° latitude to explore the sensitivity of inorganic aerosol formation to NOx emission reductions in Los Angeles
(118.239° W, 34.052° N) during COVID-19. Our detailed O3-NOx-VOC-aerosol simulations were driven by Goddard Earth
Observing System (GEOS-FP 5.22.0) assimilated meteorological fields and include anthropogenic/biogenic/biomass burning
emissions (42–44), gas-phase chemistry (45) and inorganic aerosol partitioning (46), wet/dry depositions (47–49) and transport.
We first scaled anthropogenic NOx and SO2 emissions from HTAP v2 (42) (originally for the year 2010) to the year 2017 using
satellite-derived SO2 and NOx emission reduction ratios (50) as our base emissions, which refer to emissions before lockdown
during COVID-19. We scaled our base anthropogenic NOx emissions in March by BAU/COVID monthly NOx emission ratios
from Miyazaki et al. (39) as our BAU/COVID emissions. In the COVID-19 simulations, the NOx emissions started to decrease
on March 1st.

We calculated the gas ratio (51) shown in Fig. S4 using Eq. (3):
\[
gas \text{ ratio} = \frac{[\text{NH}_3] + [\text{NH}_4^+] - 2[\text{SO}_4^{2-}]}{[\text{HNO}_3] + [\text{NO}_3^-]} \tag{3}
\]

\([\text{NH}_3], [\text{NH}_4^+], [\text{SO}_4^{2-}], [\text{HNO}_3] \text{ and } [\text{NO}_3^-] \text{ are in units of molar concentrations (mol m}^{-3}) \text{ and include both gas-phase and}
\text{aerosol-phase. This gas ratio is an indicator of NH}_4\text{NO}_3 \text{ production sensitivity to NO}_x \text{ emission change and NH}_3 \text{ emission}
\text{change. Values } > 1 \text{ indicate that NH}_4\text{NO}_3 \text{ production is NO}_x \text{ limited; values } < 1 \text{ indicate it is NH}_3 \text{ limited.}

Ozone production efficiency steady state model. The ozone production efficiency (OPE) values in Fig. S10 were computed
from a HOx-NOx steady state model similar to that used in Laughner et al. (52) (available at https://github.com/joshua-laughner/
HSSModel/releases/tag/v0.1.0, an example notebook is available at https://github.com/joshua-laughner/HOx-NOx-model-PNAS-2021).
Briefly, this model takes fixed values for NO and NO2 concentrations, VOC reactivity (VOCR), HOx productions (P(HOx)),
and RO2 + NO branching ratio (α) and solves for RO2, HO2, and OH concentrations, assuming that HO2, RO2, and the whole
HOx family (RO2 + HO2 + OH) are in steady state.

Theoretical OPE is computed from the model steady state as the ratio of ozone production to NOx loss, similar to Kleinman
et al. (53) except that formation of alkyl nitrates is counted as NOx loss:
\[
\text{OPE}_{\text{model}} = \frac{P(\text{O}_3)}{L(\text{NO}_x)} = \frac{k_{\text{NO}+\text{HO2}}[\text{NO}][\text{HO}_2] + (1-\alpha)k_{\text{NO}+\text{RO2}}[\text{NO}][\text{RO}_2]}{k_{\text{NO}_2+\text{OH}}[\text{NO}_2][\text{OH}]} \tag{4}
\]
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Table S4. Changes in NO\textsubscript{x} emissions, O\textsubscript{3}, and ozone production efficiency inferred from the multi-satellite data assimilation system.
References